

IROWG-2

Estes Park, CO, USA

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Ionospheric correction of RO signals by direct modelling of ionosphere

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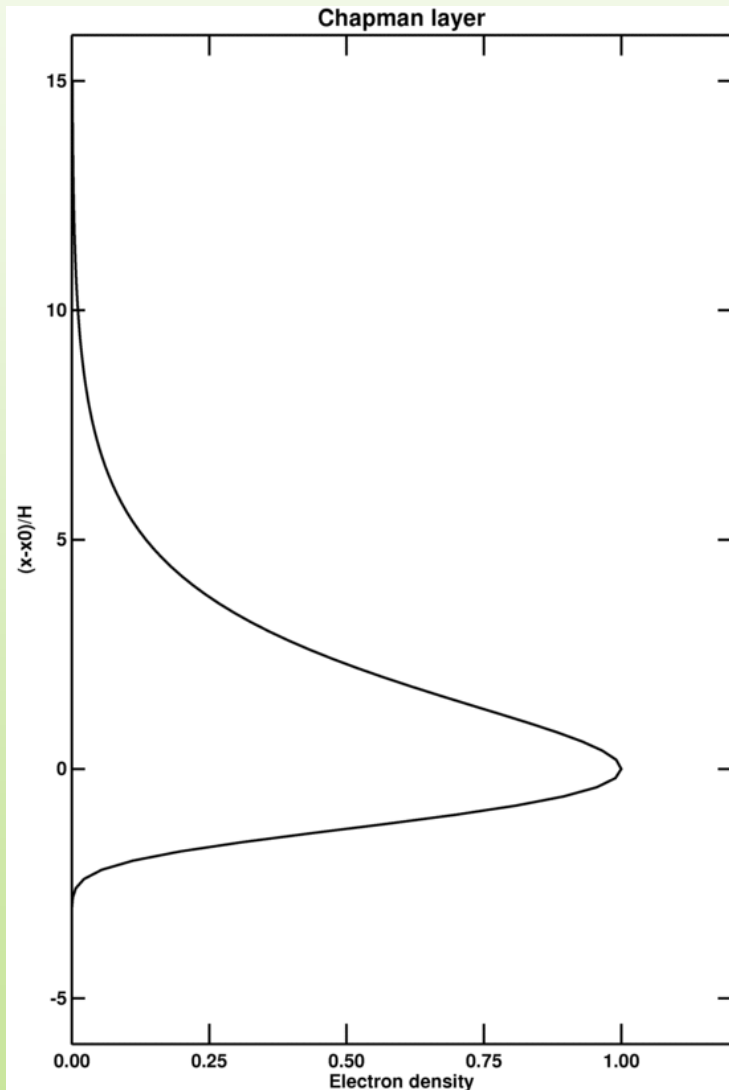


Motivation

Ultimate aim: direct assimilation of L1 and L2 signals in NWP

- How to do ionospheric correction when L2 signal drops out?
- Different (better?) tropospheric retrievals if L1 and L2 both used?
- Some information on ionospheric parameters?

Model ionosphere: electron density



Single Chapman Layer (Chapman 1931)

$$n_e(r) = \text{TEC} / \sqrt{(2\pi e H^2)} \cdot \exp\left(\frac{1}{2}(1 - u - e^{-u})\right),$$

where $u = (r - r_0) / H$.

3 parameters:

$$\text{TEC} = \int n_e dr$$

$r_0 =$ peak height

$H =$ ionospheric scale height

Model ionosphere: bending angle

$$\alpha_{Li}(a) = - 2a \int_a^\infty d \log n / dx \, dx / \sqrt{(x^2 - a^2)}, \quad x=nr$$

$$\approx (k_4 / f_{Li}^2) 2a \int_a^\infty dn_e / dx \, dx / \sqrt{(x^2 - a^2)}, \quad k_4 = 40.3 \text{ m}^3\text{s}^{-2}$$

$$\approx (k_4 / f_{Li}^2) 2a / \sqrt{(r_0+a)} \int_a^\infty dn_e / dr \, dr / \sqrt{(r - a)}$$

$$= (k_4 / f_{Li}^2) [a \text{ TEC} / \sqrt{(2\pi(r_0+a)H^3)}] \cdot Z(l)$$

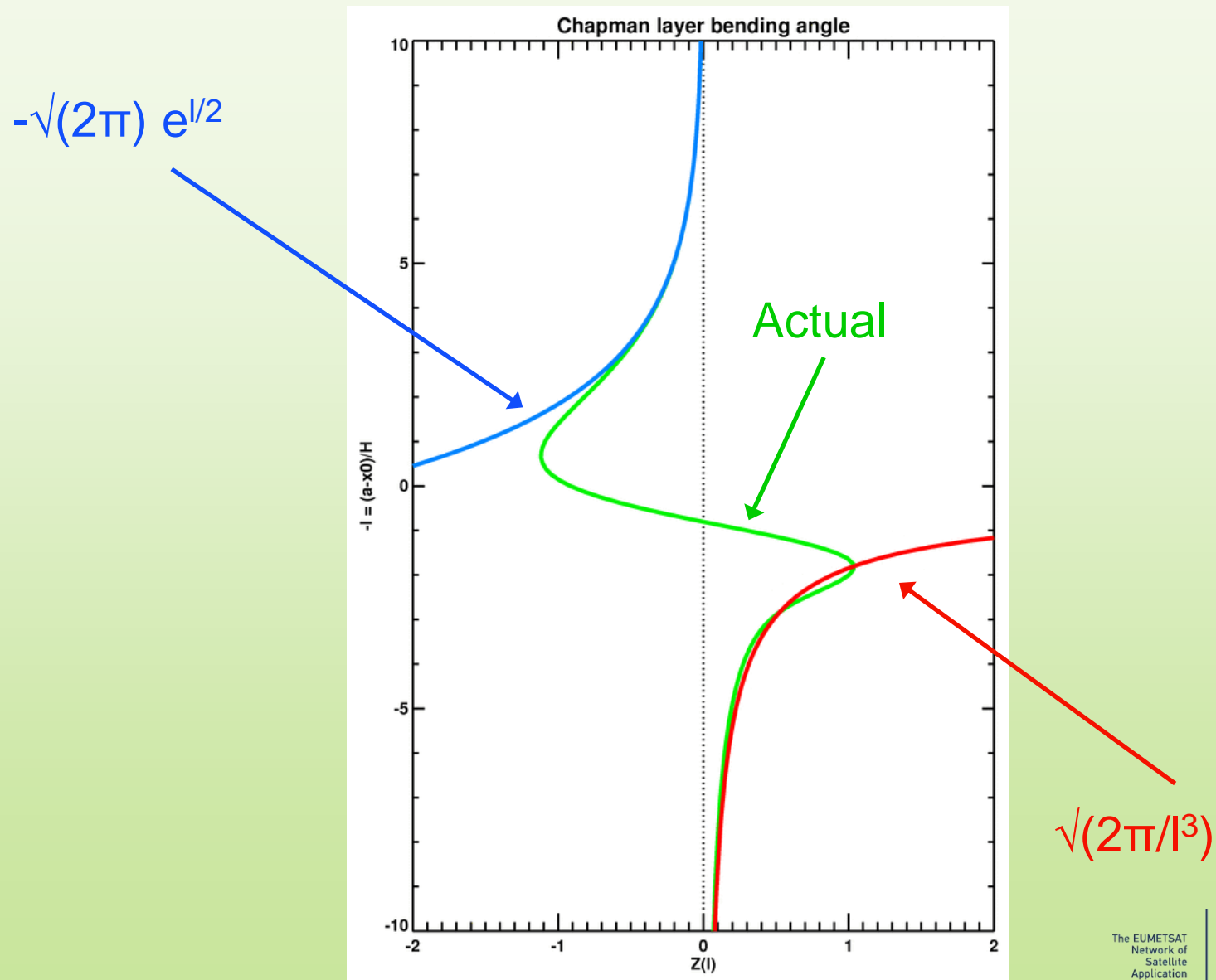
where

$$Z(l) = \int_{-l}^\infty (e^{-u}-1) \exp(1/2(1-u-e^{-u})) / \sqrt{(u+l)} \, du$$

is just a function of

$$l = (r_0 - a) / H$$

Model ionosphere: bending angle $Z(l)$

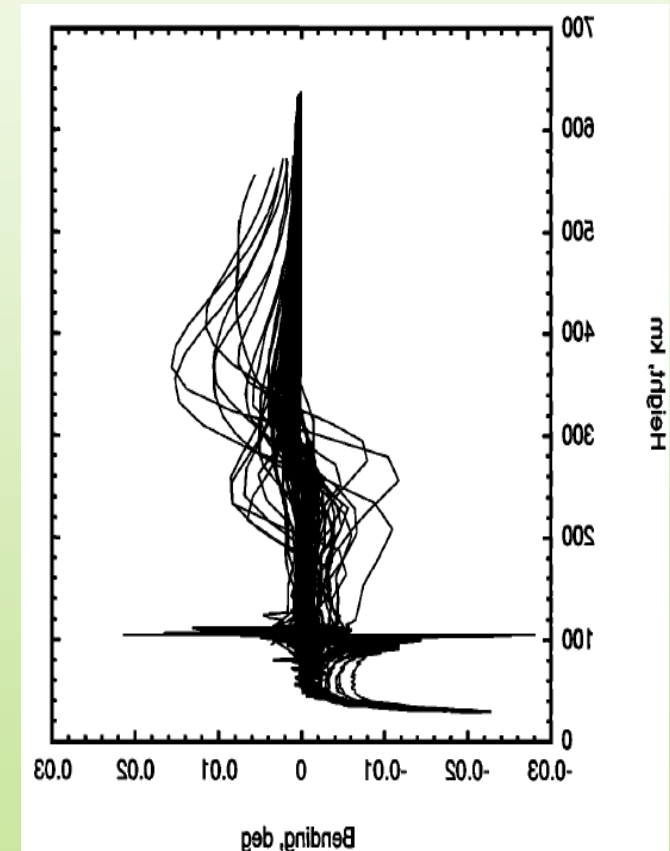
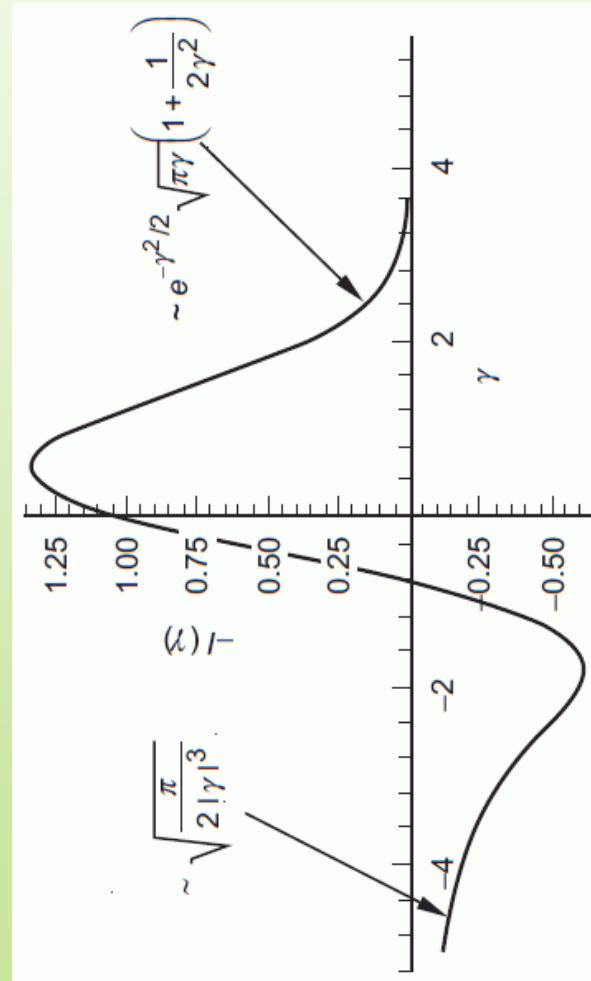
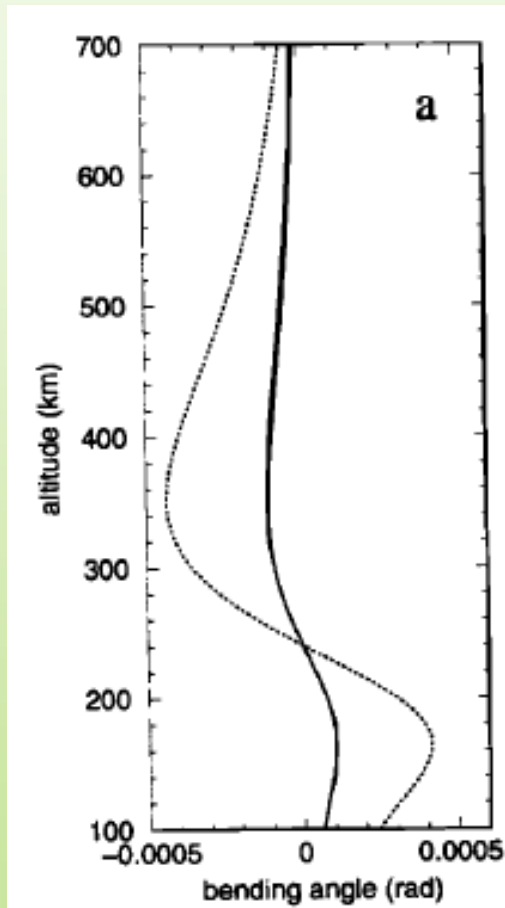


Model ionosphere: bending angle $Z(I)$

Schreiner et al, 1999,
Chapman layer

Melbourne, 2004,
Gaussian n_e

Hajj & Romans,
1998, GPS/MET



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1. Extrapolation of L2 signal

$$\alpha_{L1} = \alpha_n + k_4 f_{L1}^{-2} a \text{ TEC} / \sqrt{(2\pi(r_0+a)H^3)} Z((r_0-a)/H)$$

→

$$\alpha_n = \alpha_{L1} + (f_{L2}^2 / (f_{L1}^2 - f_{L2}^2)) \boxed{(\alpha_{L1} - \alpha_{L2})}$$

Replace

$(\alpha_{L1} - \alpha_{L2}) \sim \text{const or linear in impact parameter } a$

with

$\boxed{\text{Fitting parameter}}$

$(\alpha_{L1} - \alpha_{L2}) = \gamma a r_0 (r_0^2 - a^2)^{-3/2}$ (delta function ionosphere)

$\boxed{\text{Fitting parameter (sometimes)}}$

1. Extrapolation of L2 signal

- EUMETSAT find some weak benefit at low altitudes, but detrimental at high (stratosphere).
- Not too surprising?
- Suggests replacing

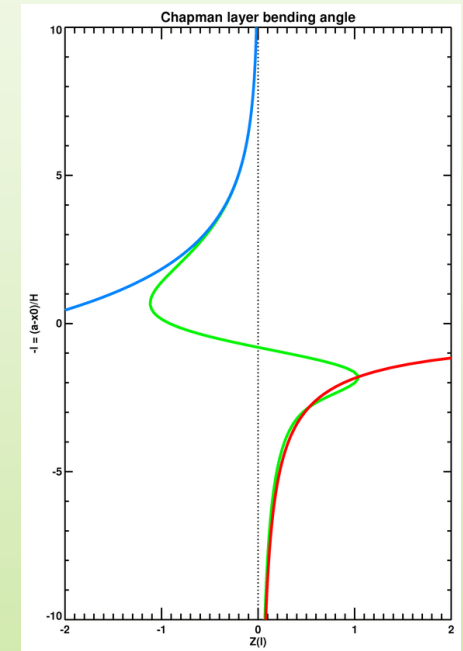
$(\alpha_{L1} - \alpha_{L2}) \sim \text{const or linear in impact parameter } a$

with

Fitting parameter

$(\alpha_{L1} - \alpha_{L2}) = \gamma a / \sqrt{(2\pi(r_0+a)H^3)} Z((r_0 - a)/H)$ (general Zorro function)

Fitting parameter (sometimes)



2. Retrievals based on L1 and L2

- Forward modelling of *model* ionosphere directly allows use of L1 and L2 separately. Is it worth it?

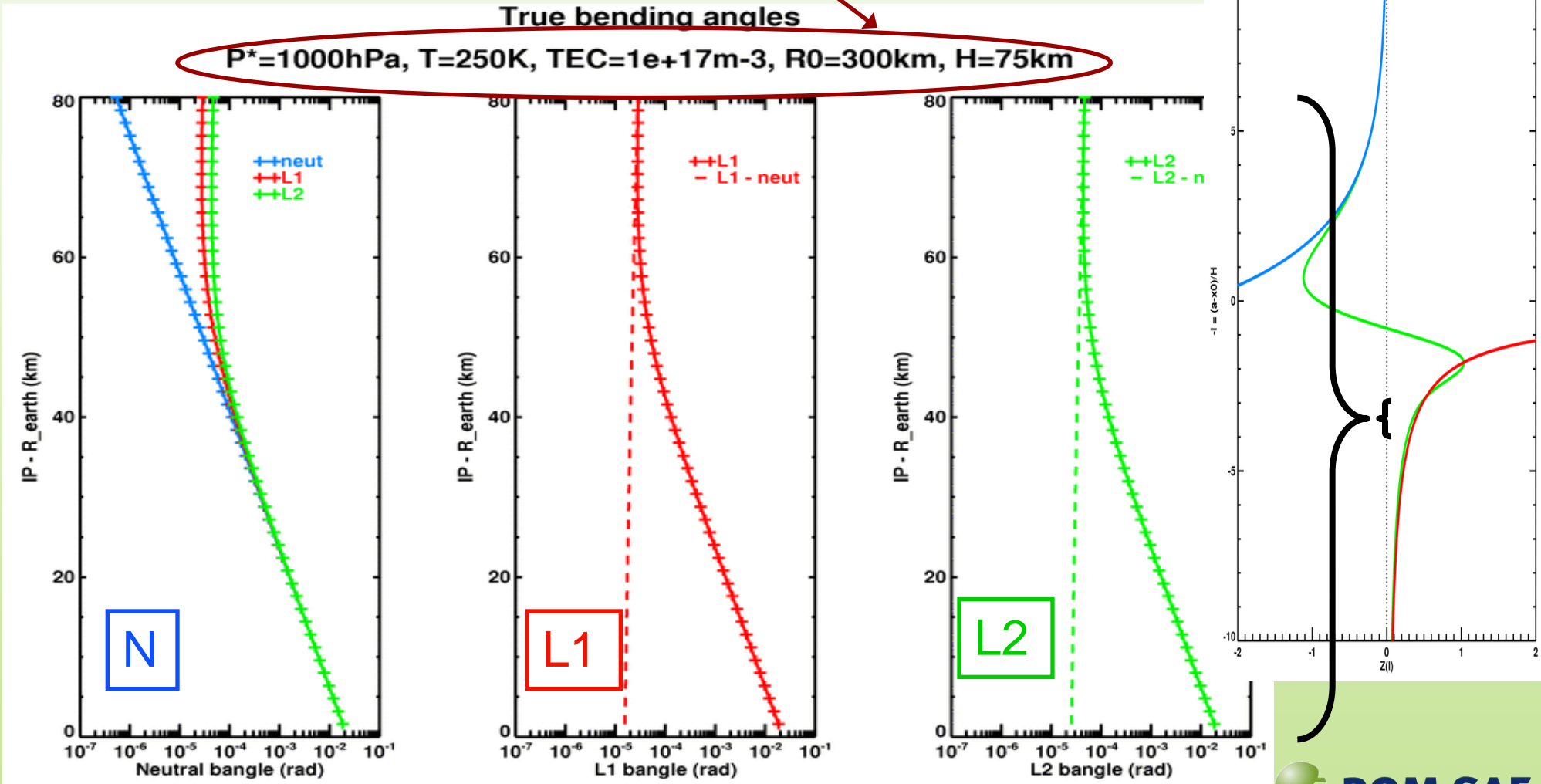
- Very simple test in a toy **0-D** model:

$$\alpha_{L1,2}(a) = k_1 (p^*/T) \sqrt{(2\pi g a / RT)} \exp(-g(a - R_E) / RT) \\ + k_4 f_{L1,2}^{-2} a \text{TEC} / \sqrt{(2\pi(r_0 + a)H^3)} Z((r_0 - a) / H)$$

- 5 element state vector: $\{T, p^*, \text{TEC}, r_0, H\}$
- Generate bending angles, add noise, perturb true state \rightarrow background, pass through minimiser, compare to “truth”.

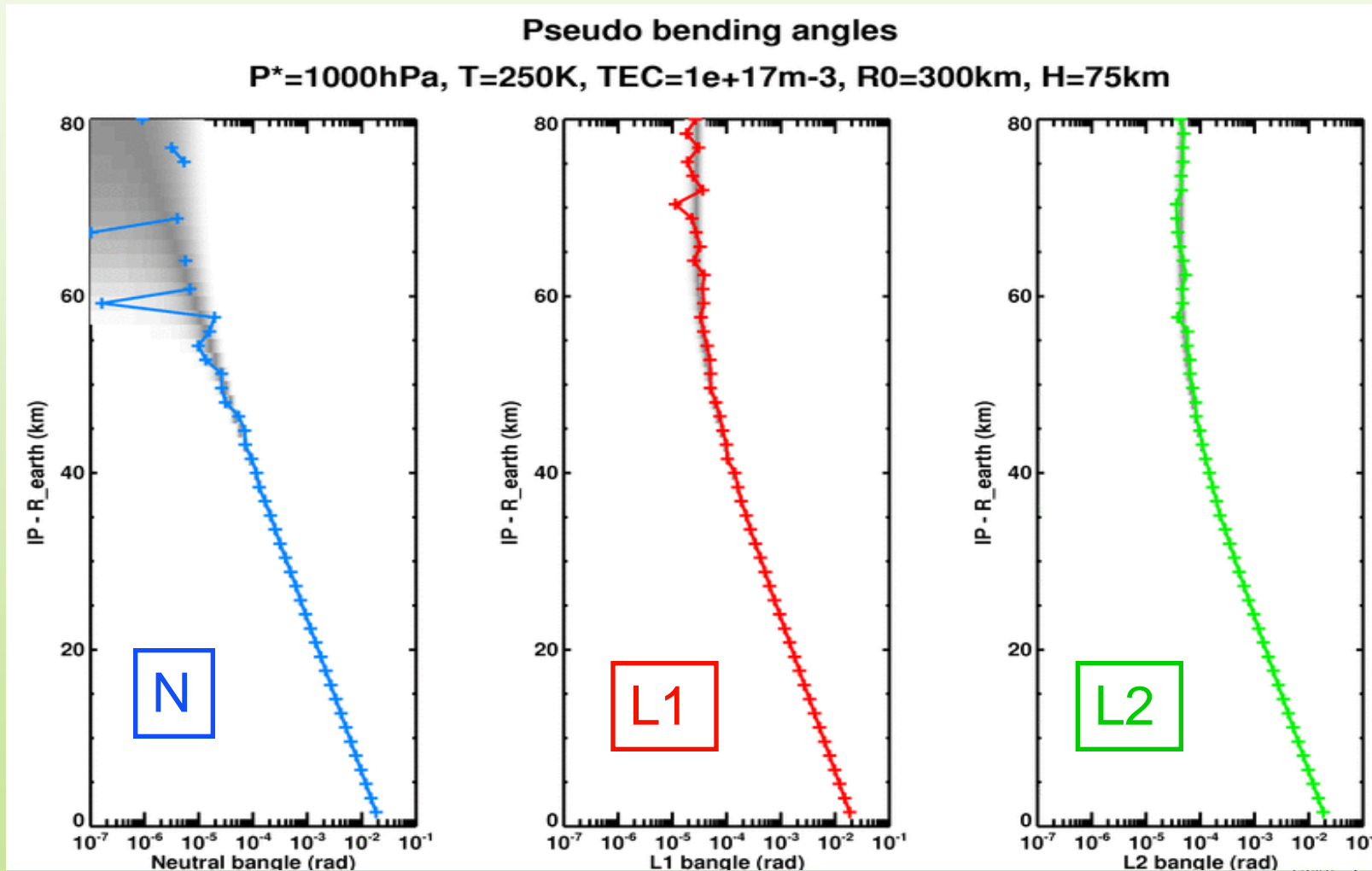
2. Retrievals based on L1 and L2

- Example 1: "standard ionosphere"



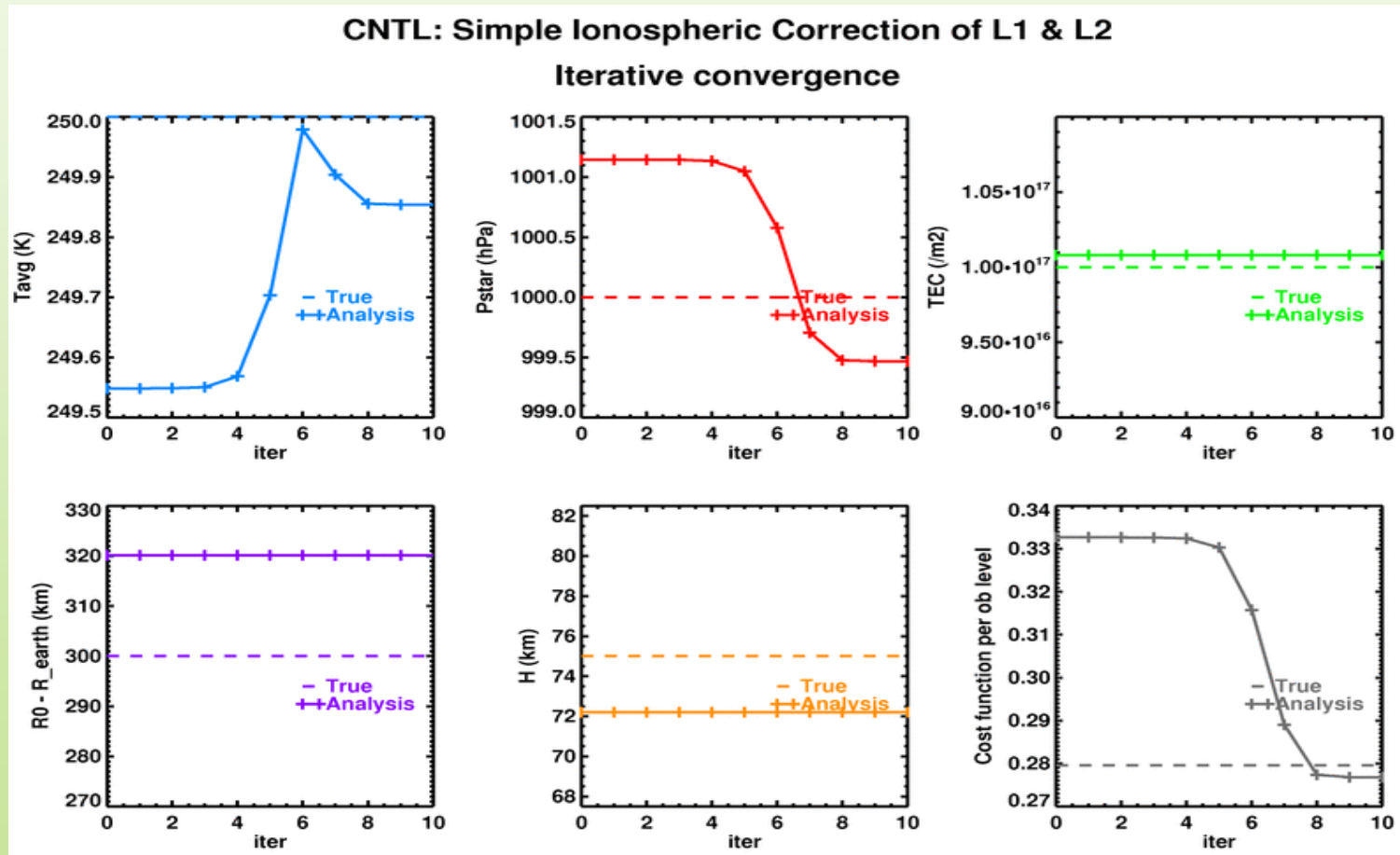
2. Retrievals based on L1 and L2

- Example 1: “standard ionosphere”



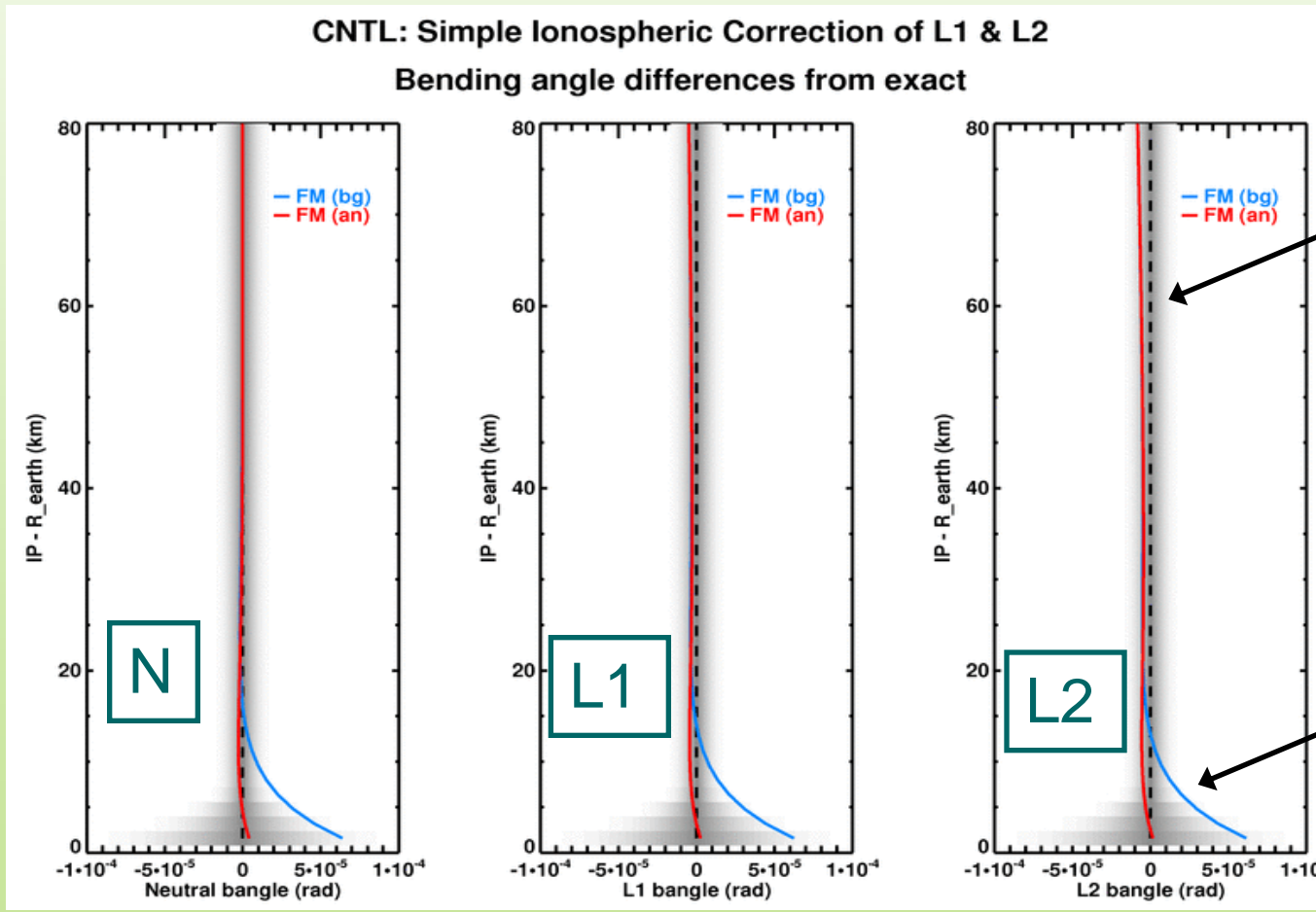
2. Retrievals based on L1 and L2

- Example 1: “standard ionosphere”.
- CNTL: 1dvar retrieval of $\{T, p^*\}$ based on $\alpha_N = ()\alpha_{L1} - ()\alpha_{L2} / () - ()$



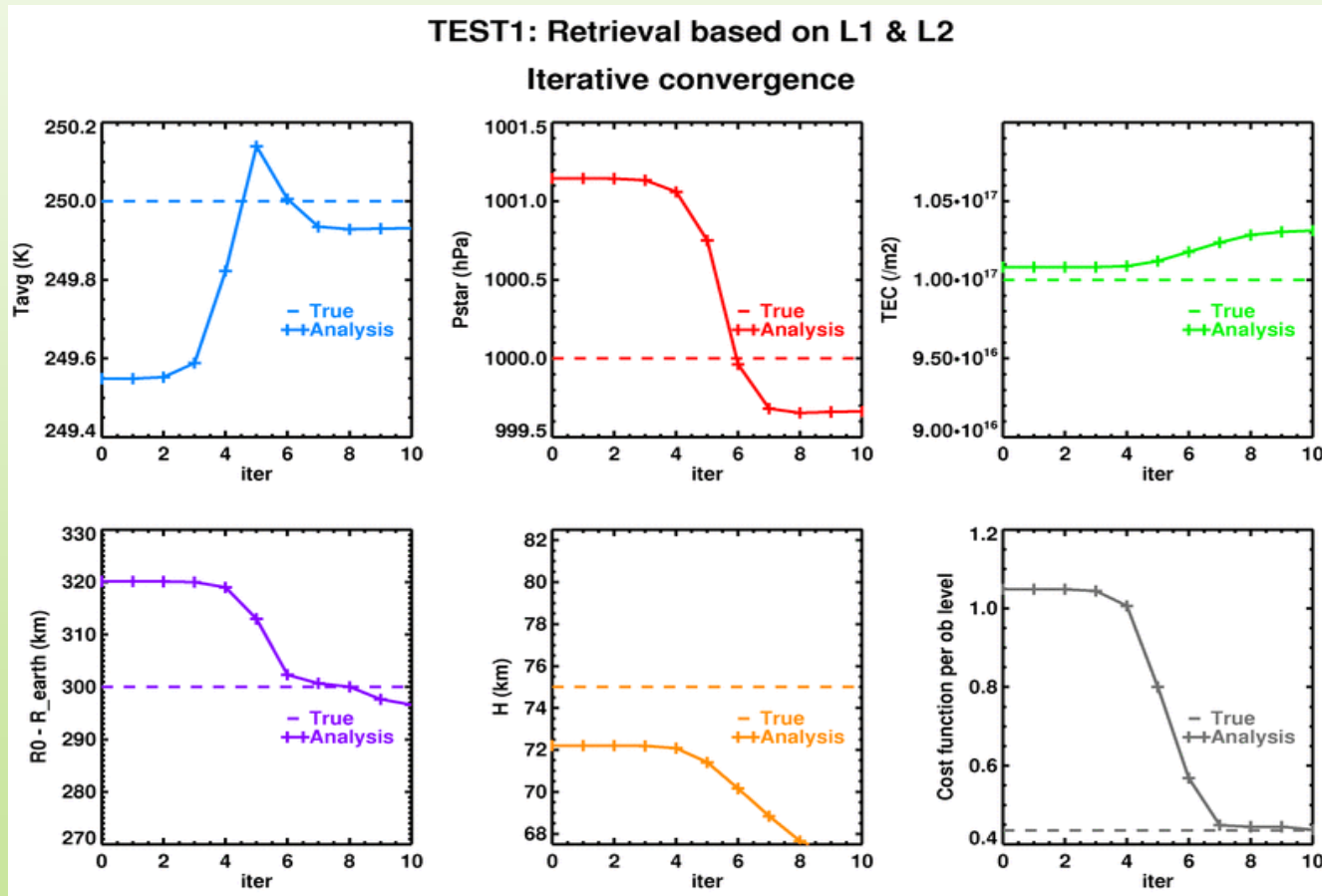
2. Retrievals based on L1 and L2

- Example 1: “standard ionosphere”.
- CNTL: 1dvar retrieval of $\{T, p^*\}$ based on $\alpha_N = (\alpha_{L1} - \alpha_{L2}) / (\alpha_{L1} - \alpha_{L2}) - (\alpha_{L2} - \alpha_{L1}) / (\alpha_{L1} - \alpha_{L2})$



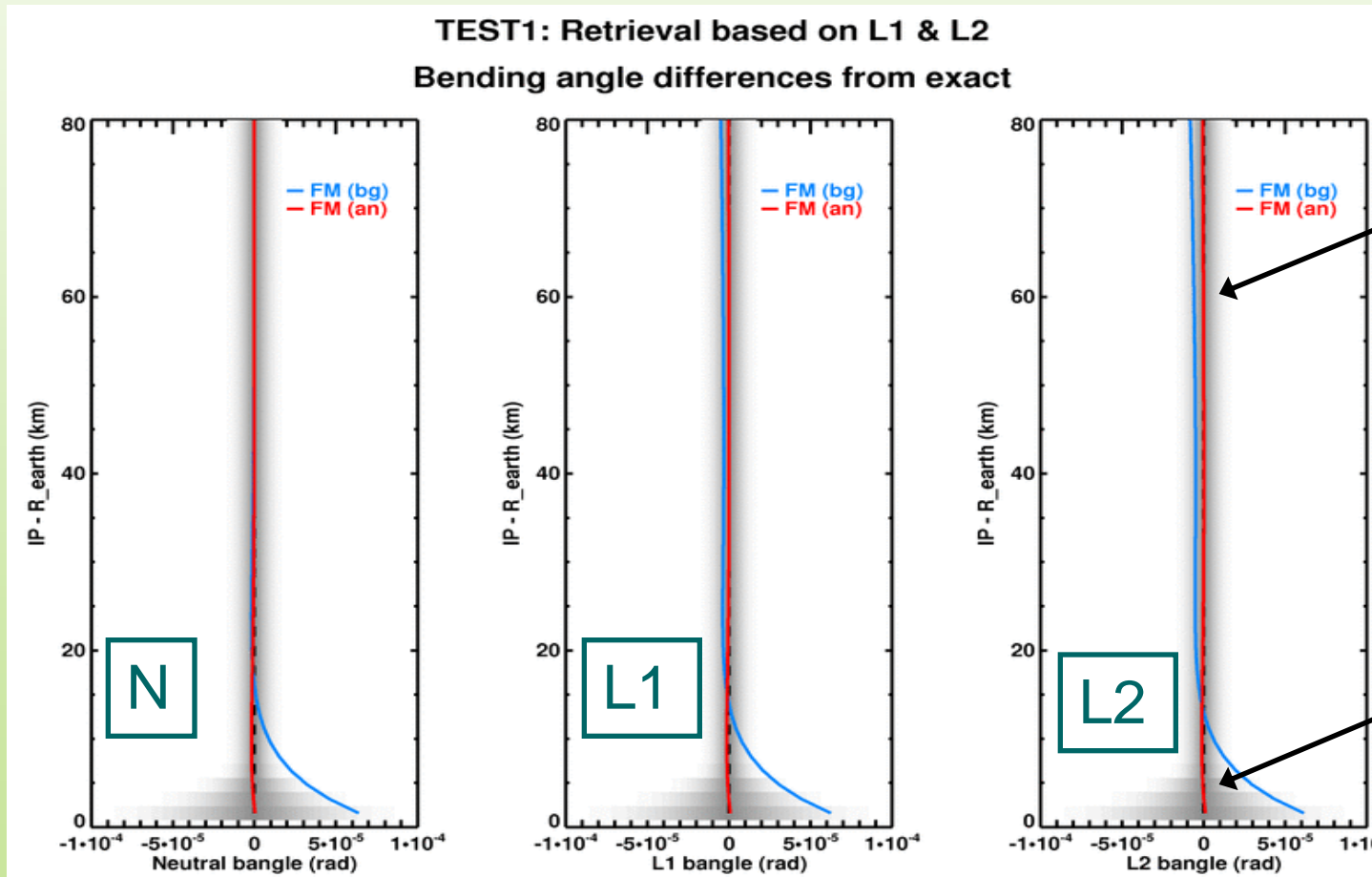
2. Retrievals based on L1 and L2

- Example 1: “standard ionosphere”.
- TEST1: 1dvar retrieval of $\{T, p^*, \text{TEC}, R_0, H\}$ based on $\{\alpha_{L1}, \alpha_{L2}\}$



2. Retrievals based on L1 and L2

- Example 1: “standard ionosphere”.
- TEST1: 1dvar retrieval of $\{T, p^*, \text{TEC}, R_0, H\}$ based on $\{\alpha_{L1}, \alpha_{L2}\}$



Better

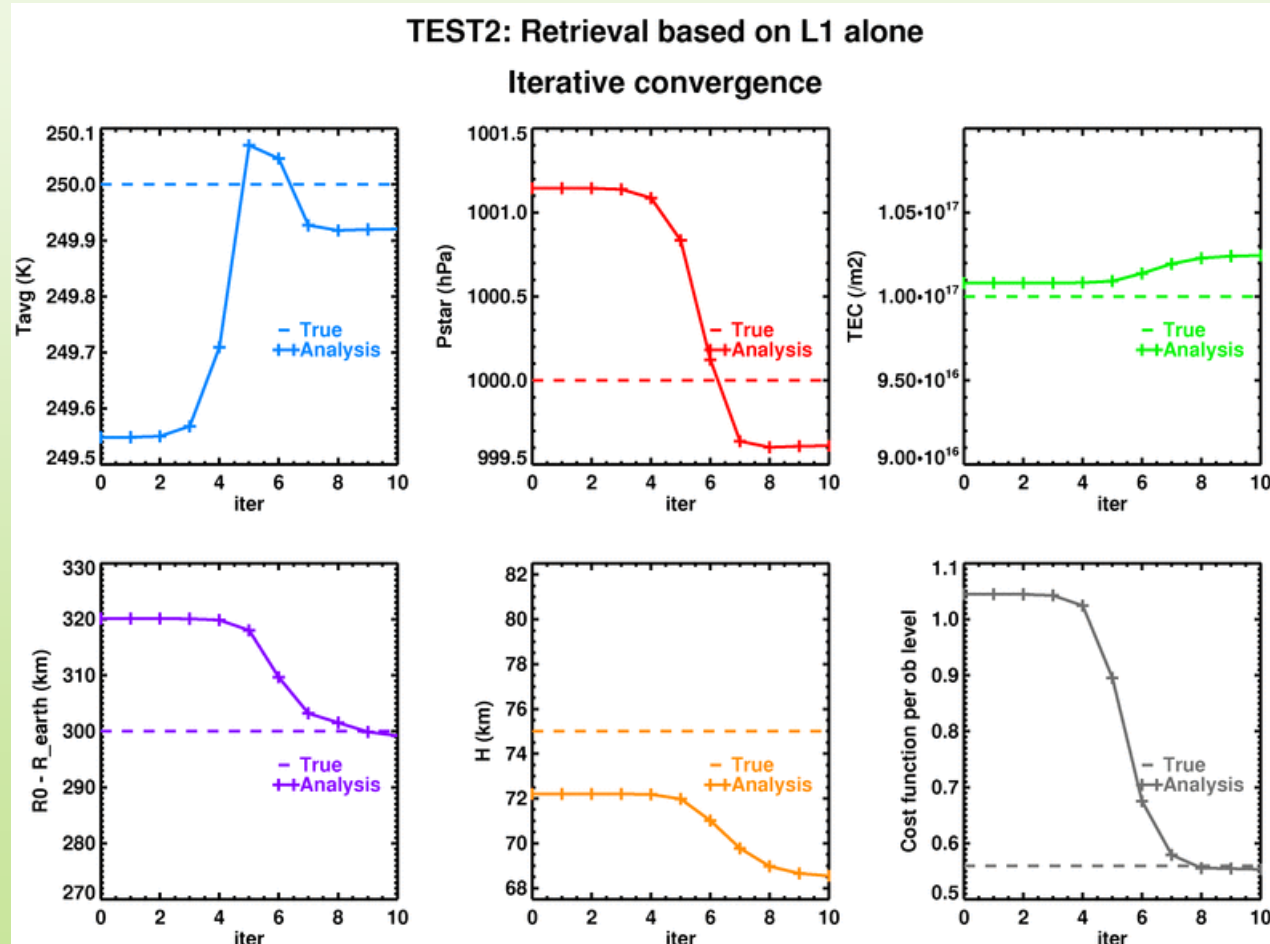
Better

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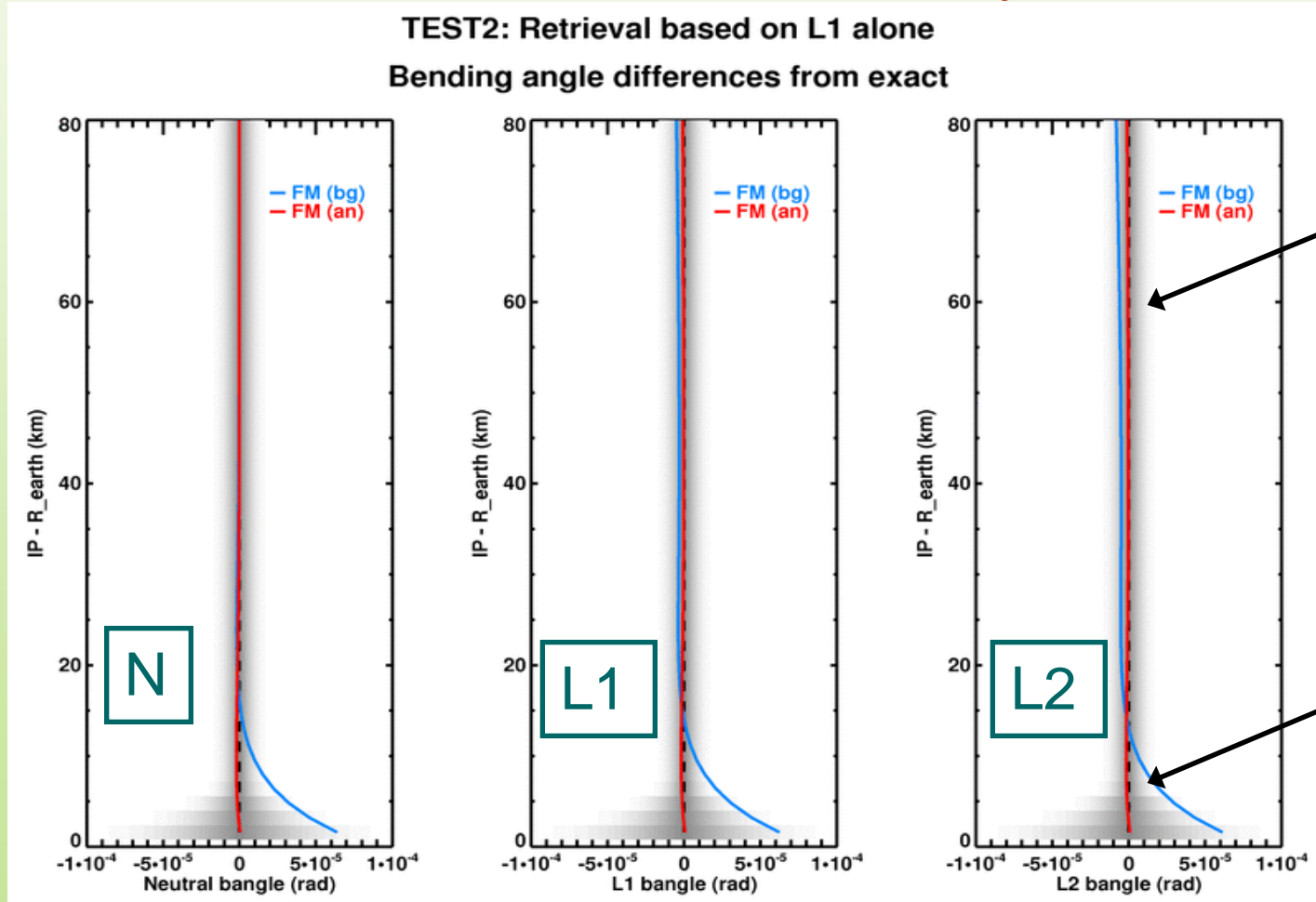
2. Retrievals based on L1 and L2

- Example 1: “standard ionosphere”.
- TEST2: 1dvar retrieval of $\{T, p^*, \text{TEC}, R_0, H\}$ based on $\{\alpha_{L1}\}$ alone



2. Retrievals based on L1 and L2

- Example 1: “standard ionosphere”.
- TEST2: 1dvar retrieval of $\{T, p^*, \text{TEC}, R_0, H\}$ based on $\{\alpha_{L1}\}$ alone



Still better

Better

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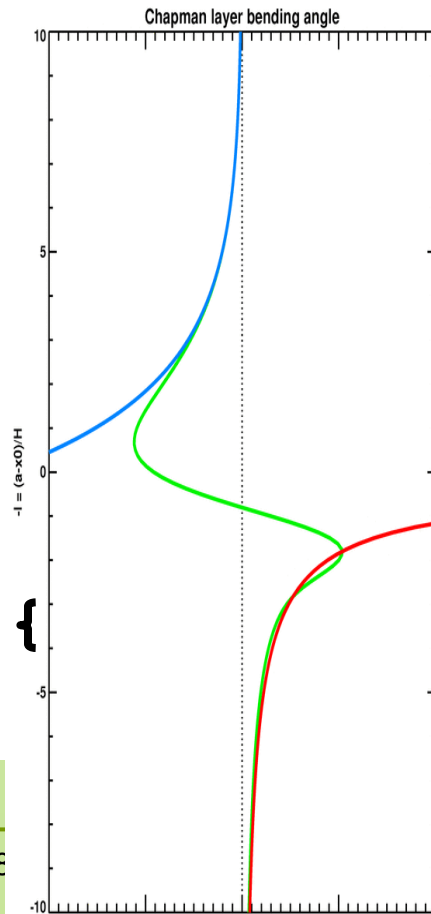
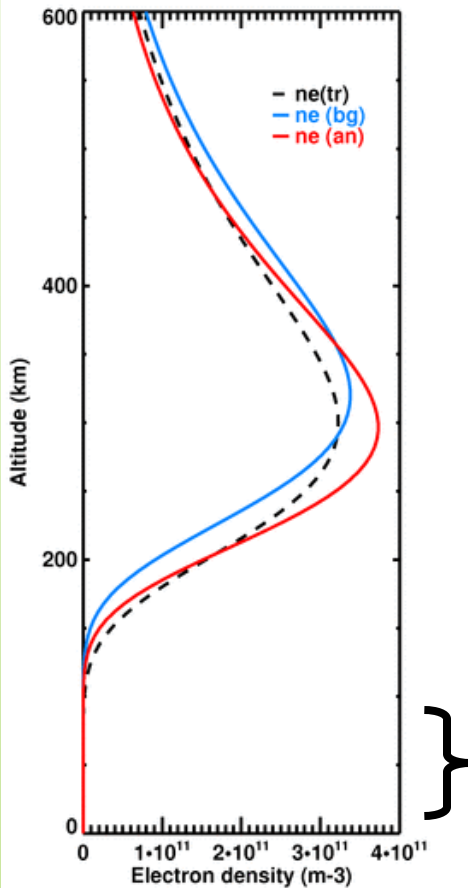
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3. Retrievals of ionospheric parameters?

- Example 1: “standard ionosphere”

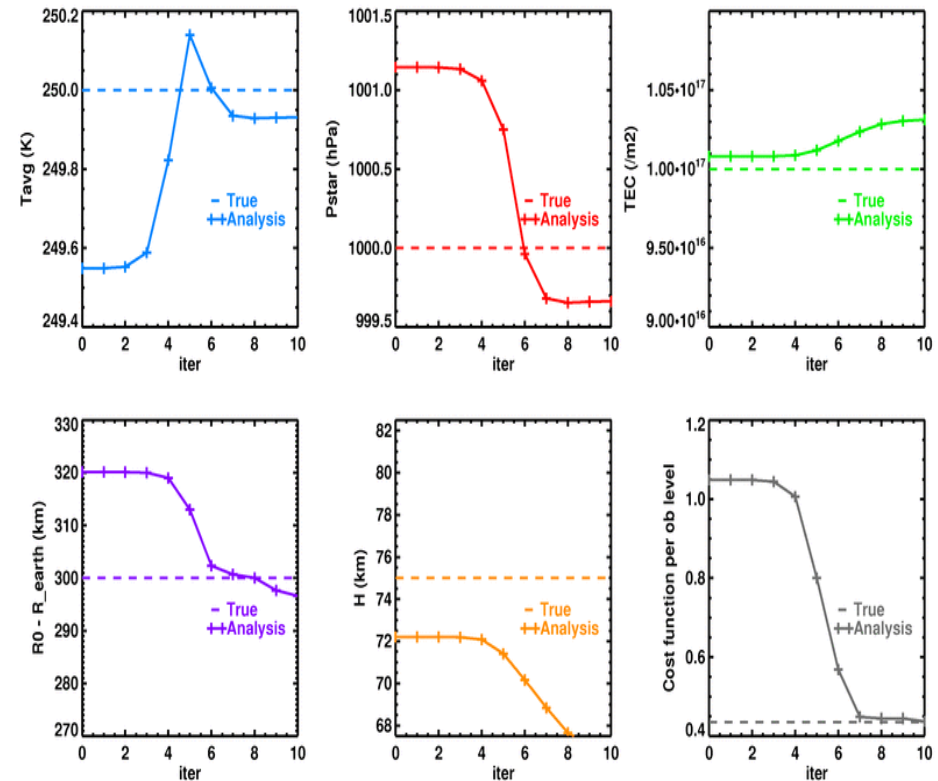
TEST1: Retrieval based on L1 & L2

$P^*=1000\text{hPa}$, $T=250\text{K}$, $\text{TEC}=1\text{e}+17\text{m}^{-3}$, $R0=300\text{km}$, $H=75\text{km}$



TEST1: Retrieval based on L1 & L2

Iterative convergence

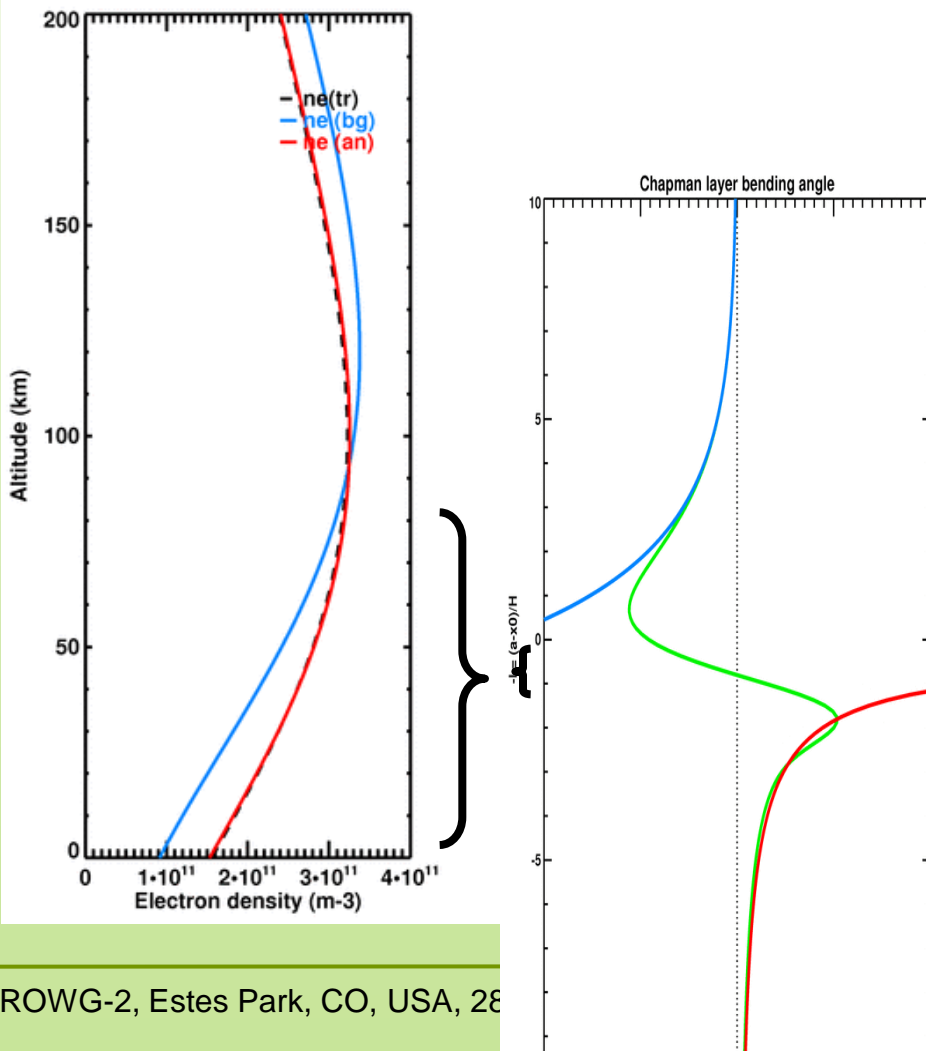


3. Retrievals of ionospheric parameters?

- Example 2: “low ionosphere”

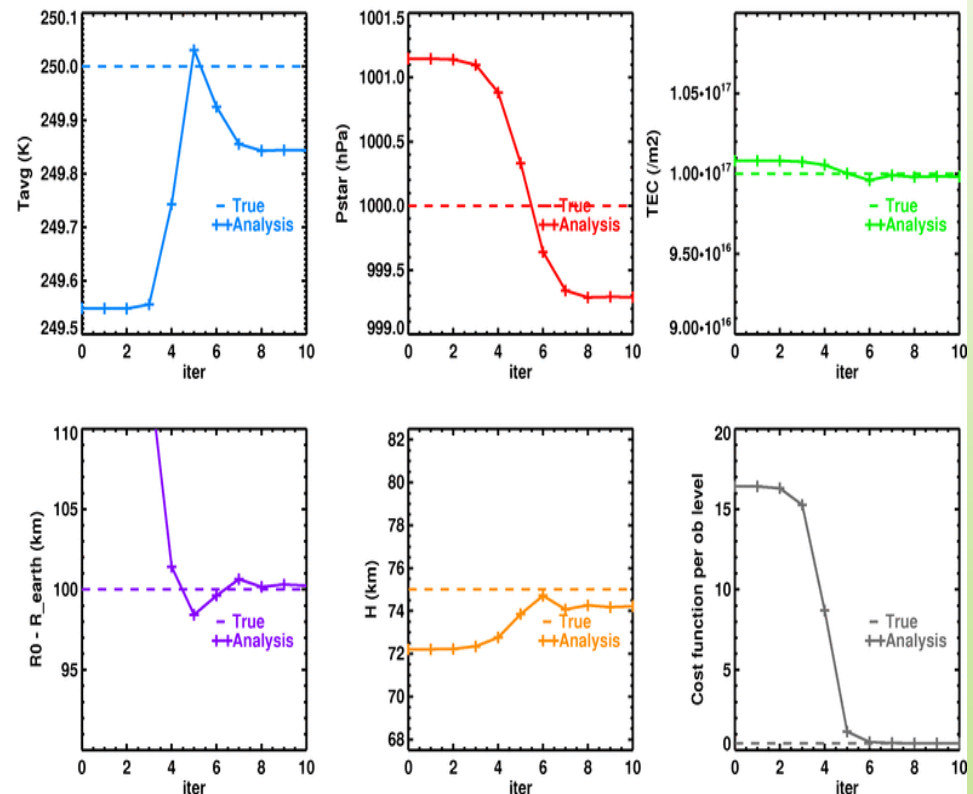
TEST1: Retrieval based on L1 & L2

$P^*=1000\text{hPa}$, $T=250\text{K}$, $\text{TEC}=1\text{e}+17\text{m}^{-3}$, $R_0=100\text{km}$, $H=75\text{km}$



TEST1: Retrieval based on L1 & L2

Iterative convergence



Conclusions

- Bending angles produced by a *model* Chapman layer ionosphere have been described
- Theory offers some hope for improved L2 extrapolation
- Might allow retrievals based on L1 and L2 signals directly (toy model)
- Probably won't allow inferences of ionospheric params to be made.